

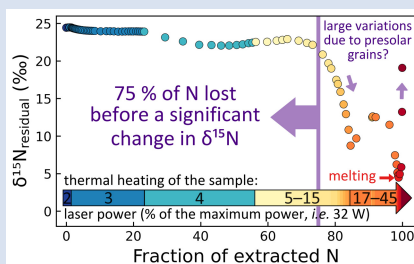
# Dissecting the complex Ne-Ar-N signature of asteroid Ryugu by step-heating analysis

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## Abstract



Samples returned from the carbonaceous asteroid (162173) Ryugu show mineralogical, chemical, and isotopic similarities with Ivuna-type (CI) carbonaceous chondrites, which likely contributed to Earth's volatile inventory. To better understand the complex Ne-Ar-N signature of CI-type material, we analysed a single, mg-sized Ryugu particle by multi-step ( $n = 85$ ) heating. Noble gases (Ne, Ar) are a mixture between implanted Solar Wind (SW), presolar component(s), and the carbonaceous phase Q, with negligible cosmogenic contributions. The  $\delta^{15}\text{N}$  variations observed during progressive heating reflect the presence of various N-bearing phases. The large number of heating steps provide key insights into the effect of thermal processing on the N abundance and isotopic ratio, and indicate that low temperatures can result in extensive N loss from CI-type material, without significantly affecting the bulk N isotopic composition. Nitrogen isotopes, therefore, remain a reliable and powerful tool for tracing volatile sources in the Solar System.

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## Introduction

The Hayabusa2 mission of the Japan Aerospace Exploration Agency (JAXA) collected 5.4 g of regolith on the Cb-type asteroid (162173) Ryugu during two touchdowns in 2019 and returned the samples to Earth on December 6<sup>th</sup>, 2020 (Tsuda *et al.*, 2020). The mineralogy and the chemical and isotopic compositions of the samples revealed a close relationship between Ryugu and Ivuna-type (CI) carbonaceous chondrites (Nakamura *et al.*, 2022; Yokoyama *et al.*, 2023). However, unlike CI chondrites collected on Earth, material from Ryugu has not been affected by heating processes during atmospheric entry or by terrestrial weathering. Since carbonaceous chondrites may have played an important role in supplying volatile elements to Earth (*e.g.*, Alexander *et al.*, 2012; Marty, 2012; Alexander, 2022), Ryugu samples are key for better understanding the origin of terrestrial volatiles. The Hayabusa2-initial-analysis volatile team and the Phase-2 curation team carried out the first noble gas and N analyses, thus providing insights into the volatile composition, formation, and alteration history of Ryugu (Nakamura *et al.*, 2022; Okazaki *et al.*, 2023). The noble gases were found to be mainly primordial (*i.e.* carried by the so called phase Q and a variety of presolar components), with variable contributions from solar wind (SW) and cosmogenic isotopes. Okazaki *et al.* (2023) concluded that the heterogeneous N contents and isotopic compositions of Ryugu samples, including four small pelletised samples and three splits of a large aggregate sample (Naraoka *et al.*, 2023), indicate the presence of at least two carrier phases: a N-rich phase with  $\delta^{15}\text{N}$  up to +70 ‰ and a N-depleted phase with  $\delta^{15}\text{N}$  near 0 ‰. These results were subsequently discussed in more detail by Broadley *et al.* (2023) and Hashizume *et al.* (2024).

Here, we aim to better understand the nature and behaviour of Ne-Ar-N carriers in Ryugu material by performing a large number ( $n = 85$ ) of extraction steps at incrementally increasing temperature. This is the first time, to our knowledge, that a single asteroidal particle has been heated in so many steps for coupled noble gas and N analyses, thereby providing unprecedented insight into the complex Ne-Ar-N makeup of primitive extraterrestrial matter.

## Material and Methods

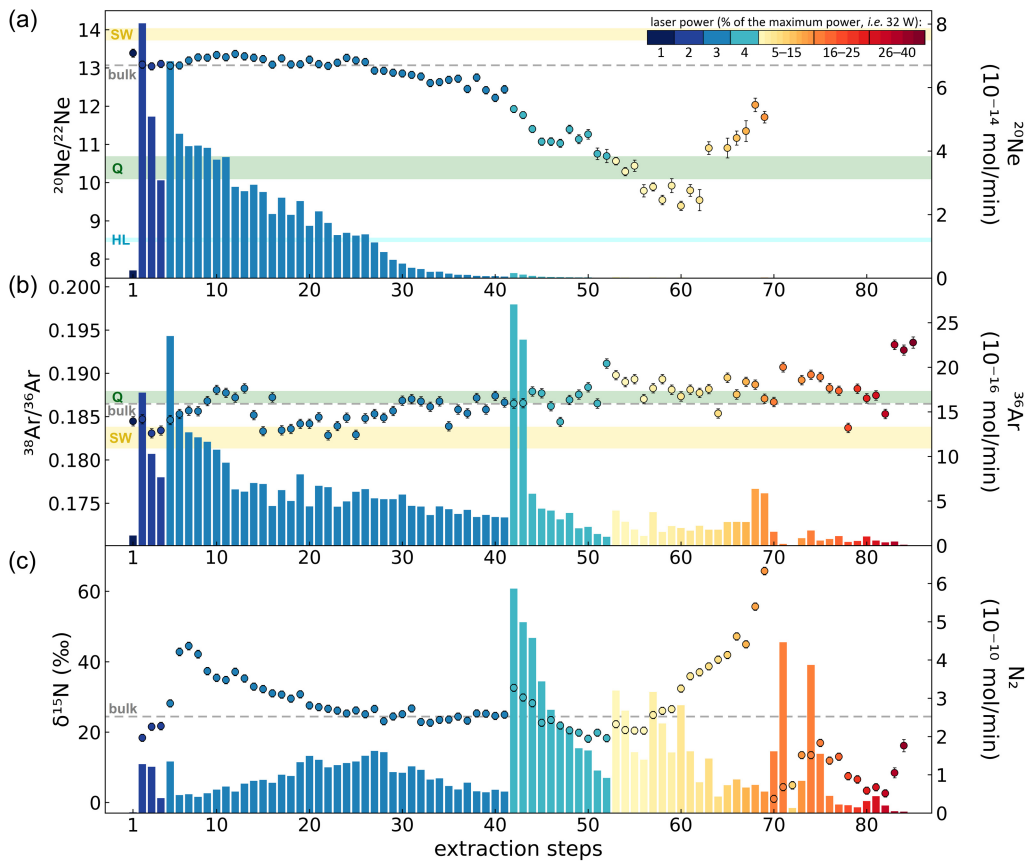
Particle C0015 ( $1.8 \pm 0.2$  mg), collected during the second touchdown on Ryugu within a crater created by the spacecraft's small carry-on impactor, was targeted for step-heating Ne-Ar-N analysis at the Centre de Recherches Pétrographiques et Géo-chimiques (CRPG) noble gas facility. The particle was never in contact with Earth's atmosphere and was constantly kept within a dry  $\text{N}_2$  atmosphere or under vacuum during sample preparation, storage, and analysis. The particle was heated under static vacuum using a  $\text{CO}_2$  laser at increasing laser power, and the fraction of noble gases (Ne, Ar) and N (in the form of  $\text{N}_2$ ) extracted at each step was analysed using a Noblesse-HR noble gas mass spectrometer in multi-collection. A total of 85 heating steps were performed to study the progressive release of different noble gas and N components. Details on the analytical procedure and data treatment are provided in the Supplementary Information.

## Ne-Ar-N Release Patterns of Particle C0015

Figure 1 shows the release patterns of Ne, Ar, and N (analysed in the form of  $\text{N}_2$ ) and the corresponding isotope ratios measured

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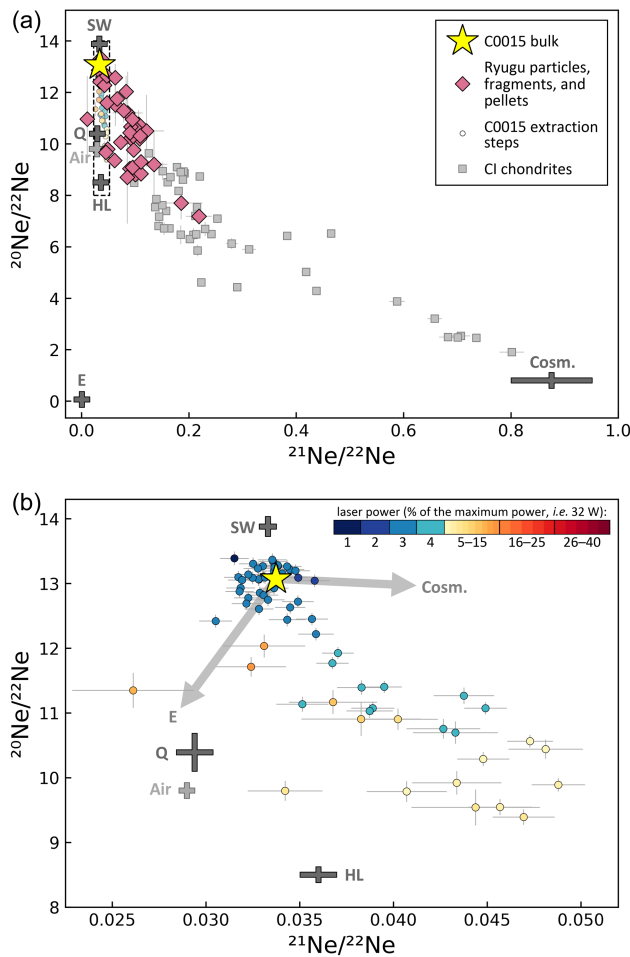
**Figure 1** (a)  $^{20}\text{Ne}$ , (b)  $^{36}\text{Ar}$ , and (c)  $\text{N}_2$  abundances (filled bars) and isotope ratios (filled circles) of particle C0015 for 85 extraction steps. Abundances are weighted by the extraction duration (*i.e.* 4 min for the first 26 steps and 12 min for the following 59 steps). Isotope ratios are shown for heating steps with  $<30\%$  blank contributions. The different colours represent the laser power used for the extraction (*i.e.* 1 to 45 % of the total power of 32 W). The dashed lines indicate the calculated “bulk”  $^{20}\text{Ne}/^{22}\text{Ne}$ ,  $^{38}\text{Ar}/^{36}\text{Ar}$ , and  $\delta^{15}\text{N}$  values. Isotope ratios of the solar wind (SW), phase Q, and Ne-HL are shown for comparison (see Ott, 2014 and references therein). Uncertainties are  $1\sigma$  and error bars are, in most cases, smaller than symbol sizes.

at each heating step. Most of the Ne was released at very low laser power ( $\leq 3\%$ ) at which the camera view showed no visible heating (indicated by an orange glow) of the particle. The  $^{20}\text{Ne}/^{22}\text{Ne}$  ratio was nearly constant ( $12.92 \pm 0.47$ ) during the first 38 extraction steps (Fig. 1a), only slightly below the Ne isotope ratio of SW ( $13.74 \pm 0.02$  to  $14.00 \pm 0.04$ ; Ott, 2014) and close to that of the protosolar nebula ( $\sim 13.36$ ; Heber *et al.*, 2012). This plateau likely results from the extraction of a pure Ne component. While trapping of nebular gas in CI-type material cannot be ruled out, the  $^3\text{He}/^4\text{He}$  ratio of solar-gas-rich Ryugu samples analysed previously is consistent with implanted SW-derived gas (Meshik *et al.*, 2023; Okazaki *et al.*, 2023). Neon released at low temperatures must, therefore, be predominantly derived from the SW, fractionated to an isotopically lighter value upon implantation and grain surface sputtering (Grimberg *et al.*, 2006). At higher temperatures, the  $^{20}\text{Ne}/^{22}\text{Ne}$  ratio first decreased significantly to  $9.39 \pm 0.12$ , before increasing to  $12.03 \pm 0.18$ . Since the  $^{21}\text{Ne}/^{22}\text{Ne}$  ratio varied only between  $\sim 0.026$  and  $0.049$  (Fig. 2), which implies a small proportion of cosmogenic Ne (*i.e.*  $<0.1\%$   $^{21}\text{Ne}_{\text{cosm}}$ ) in particle C0015, the  $^{20}\text{Ne}/^{22}\text{Ne}$  variations are inferred to predominantly reflect the release of the primordial components Ne-Q and Ne-HL. The Ne amounts released during the last 15 steps were too low for reliable isotope ratio measurements. The results indicate the presence of at least three Ne components in Ryugu particle C0015: SW-derived Ne implanted at the grain surface, Ne-HL carried by presolar nanodiamonds, and Ne-Q (or P1) carried by phase Q (Fig. 2b). Overall, the Ne composition of particle

C0015 is dominated by the SW component (Fig. 2a), similar to two pellets studied by the Hayabusa2-initial-analysis volatile team (Okazaki *et al.*, 2023) and several particles analysed by the Phase-2 curation team (Nakamura *et al.*, 2022).

The  $^{40}\text{Ar}$  signal was small for steps #2 and #5, and comparable to the blank value throughout the rest of the heating procedure, demonstrating the absence of any adsorbed atmospheric Ar. The  $^{38}\text{Ar}/^{36}\text{Ar}$  ratio of the Ar fraction released at low laser power first oscillated between the Ar-SW ( $0.1818 \pm 0.0005$  to  $0.1828 \pm 0.0010$ ; Ott, 2014) and Ar-Q values ( $0.1872 \pm 0.0007$ ; Wieler *et al.*, 1992), then plateaued around Ar-Q (Fig. 1b). Ar-HL ( $0.227 \pm 0.003$ ; Ott, 2014), expected to be released from presolar diamonds at relatively low temperatures (Supplementary Information), likely also contributed to the observed variations. Unlike Ne, a large amount of Ar was released at higher temperatures, implying that a significant proportion of Ar in particle C0015 was carried by refractory phases. The elevated  $^{38}\text{Ar}/^{36}\text{Ar}$  ratios observed during sample melting (*i.e.* after step #80, at  $\geq 24\%$  laser power) can be explained by a very small contribution of cosmogenic  $^{38}\text{Ar}$  ( $^{38}\text{Ar}/^{36}\text{Ar}_{\text{cosm}} \sim 1.54$ ; Wieler, 2002). The calculated bulk  $^{38}\text{Ar}/^{36}\text{Ar}$  ratio ( $0.1865 \pm 0.0001$ ; Table S-1) is comparable to that of Ar-Q and similar to the values previously reported for Ryugu ( $0.186 \pm 0.001$  to  $0.194 \pm 0.007$ ; Okazaki *et al.*, 2023).

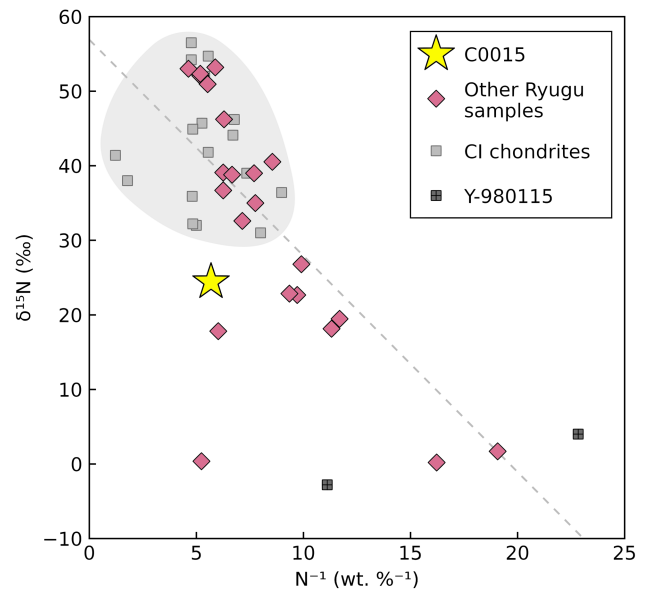
The amount of N released and the corresponding  $\delta^{15}\text{N}$  varied significantly for the different extraction steps, reflecting the presence of several, isotopically distinct N components in Ryugu material. The blank contribution exceeded 30 % of the



**Figure 2** (a) Neon isotopic composition of Ryugu particle C0015 (bulk represented by the yellow star; individual extraction steps represented by the small circles) compared to pelletised and fragment Ryugu samples analysed by the Hayabusa2-initial-analysis volatile team and particles analysed by the Phase-2 curation team (Institute for Planetary Materials, Okayama University) (pink diamonds), as well as CI chondrites (grey squares) (Nakamura *et al.*, 2022; Broadley *et al.*, 2023; Meshik *et al.*, 2023; Okazaki *et al.*, 2023 and references therein). Different Ne end members are represented by grey crosses (Ott, 2014). Ne-E is mainly carried by presolar graphite and silicon carbide. (b) Close-up (corresponding to the dashed rectangle in panel (a)) of the Ne isotopic variations observed for the 85 extraction steps of particle C0015. The different colours represent the laser power used for each extraction. Uncertainties are 1σ.

measured N<sub>2</sub> signal at step #1, and a δ<sup>15</sup>N of +18.4 ± 1.0 ‰ was measured at step #2. This indicates that the contribution of any dry N<sub>2</sub> adsorbed onto the sample surface during sample storage and handling must be negligible, provided that this component is characterised by an atmospheric isotope signature. Whereas the δ<sup>15</sup>N ranged from +18.1 ± 1.0 up to +44.5 ± 1.1 ‰ during the first 52 steps (*i.e.* at a low laser power, ≤4 %), δ<sup>15</sup>N increased to +65.8 ± 1.1 ‰ at step #69 and then suddenly dropped to +1.0 ± 1.0 ‰ at step #70. No clear SW contribution (δ<sup>15</sup>N<sub>SW</sub> = -407 ± 7 ‰; Marty *et al.*, 2011) could be identified, even for low temperature extraction steps. Thus, the N release pattern is clearly decoupled from that of Ne. Assuming an unfractionated SW <sup>20</sup>Ne/<sup>14</sup>N elemental abundance ratio of 1.14 (Marty *et al.*, 2010), any SW-derived N is indeed expected to be undetectable in the <sup>20</sup>Ne-rich particle C0015.

The production of cosmogenic <sup>15</sup>N during exposure to cosmic rays can potentially modify the N isotope ratio of



**Figure 3** Nitrogen isotopic composition (δ<sup>15</sup>N) as a function of the inverse of the N concentration of Ryugu samples compared to CI chondrites and Y-980115, a CI chondrite recovered in Antarctica that has recently been recommended to be re-classified as a Yamato-type (CY) carbonaceous chondrite (King *et al.*, 2019). Adapted from Okazaki *et al.* (2023) and Hashizume *et al.* (2024), including data reported in Table S-2. A simple two-component mixture, between a N-rich phase with δ<sup>15</sup>N up to +70 ‰ and a N-depleted phase with δ<sup>15</sup>N near 0 ‰, fails to explain the N signature of Ryugu particle C0015.

extraterrestrial samples. Combined analyses of noble gases and N make it possible to quantify the amount of <sup>15</sup>N<sub>cosm</sub> and to identify the isotopic composition of primordial N. Previous Ne analyses revealed that Ryugu samples record short cosmic ray exposure (CRE) ages of ~100 kyr to 8 Myr (Nakamura *et al.*, 2022; Okazaki *et al.*, 2023). The Ne isotopic composition of particle C0015 studied here closely resembles that of pellets A0105-15 and A0105-06, which are dominated by Ne-SW and contain a negligible amount of cosmogenic Ne. By using the same <sup>21</sup>Ne<sub>cosm</sub> production rate as Okazaki *et al.* (2023) (*i.e.* P<sub>21</sub> = 1.34 × 10<sup>-13</sup> mol g<sup>-1</sup> Myr<sup>-1</sup>), together with a simplified one-stage irradiation model, the CRE age of particle C0015 is estimated at ~30,000 years. This exposure duration could result in the production of ~4 × 10<sup>-17</sup> mol <sup>15</sup>N (when using the maximum (<sup>15</sup>N/<sup>21</sup>Ne)<sub>cosm</sub> = 5.5 from Mathew and Murty, 1993, estimated for H/L chondrites), which is negligible compared to the bulk N content of the sample. Since particle C0015 has neither been measurably affected by the production of cosmogenic <sup>15</sup>N nor atmospheric contamination and N-SW implantation, it preserves key information on N components and carrier phases in Ryugu and CI-type material. However, distinguishing between different labile and more refractory N components in particle C0015 is challenging because, in contrast to the noble gases, the nature and N isotope compositions of potential chondritic N carrier phases are highly complex, as detailed in the Supplementary Information. While the different N carrier phases cannot be deciphered, the observed N release pattern (Fig. 1c) and the summed bulk N signature (Fig. 3) require the presence of more than just two N components, in contrast to previous findings of Okazaki *et al.* (2023) and Hashizume *et al.* (2024).

## Discussion

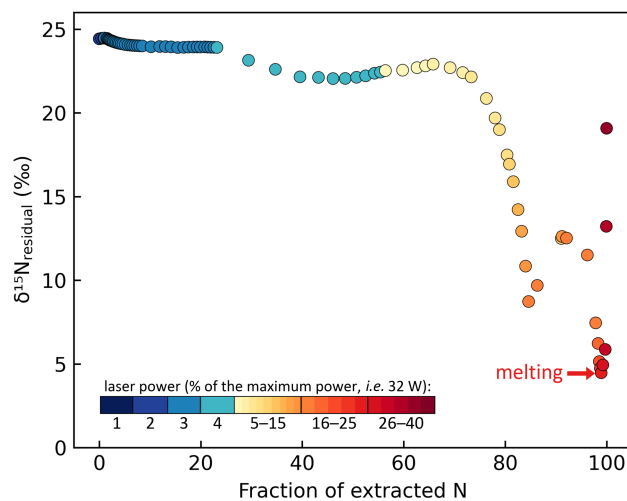
The mineralogical, chemical, and isotopic characteristics of Ryugu revealed a close relationship with CI-chondrites

(Nakamura *et al.*, 2022; Yokoyama *et al.*, 2023). The CI group nominally comprises the five falls Orgueil, Alais, Ivuna, Tonk, and Revelstoke, whose bulk N isotopic compositions vary between +31 and +61 ‰ (based on analyses of Alais, Ivuna, and Orgueil; Table S-2 and references therein) (Fig. 3). It is noteworthy that the Meteoritical Bulletin lists an additional four CIs that were recovered in the Yamato Mountains in Antarctica (Y-86029, Y-86737, Y-980115, and Y-980134), two of which (Y-86029 and Y-980115) have been recommended to be re-classified as Yamato-type (CY) chondrites (King *et al.*, 2019). The bulk  $\delta^{15}\text{N}$  of Y-980115 was reported to vary from  $-2.8$  ‰ to  $+4.0$  ‰ (Chan *et al.*, 2016; Hashizume *et al.*, 2024) (Fig. 3), distinct from the CI signature.

The N abundances and isotopic compositions of several particles collected at Ryugu's surface (Ryugu-A samples) or within an artificial crater (Ryugu-C samples) fall within the CI range; however, some samples show lower N contents and/or  $\delta^{15}\text{N}$  values (Fig. 3; Nakamura *et al.*, 2022; Broadley *et al.*, 2023; Naraoka *et al.*, 2023; Oba *et al.*, 2023; Hashizume *et al.*, 2024). The four pelleted samples analysed by the Hayabusa2-initial-analysis volatile team have particularly low N contents (Table S-2). The bulk  $\delta^{15}\text{N}$  of  $+24.43 \pm 0.17$  ‰ of particle C0015 (Table S-1) is comparable to the values previously obtained for the pellets A0105-05 and C0106-06 by noble gas mass spectrometry at CRPG ( $+18.14 \pm 0.94$  ‰ and  $+19.47 \pm 0.89$  ‰; Broadley *et al.*, 2023; Okazaki *et al.*, 2023), whereas its summed bulk N abundance ( $1760 \pm 195$  ppm; Table S-1) is consistent with the range observed in most CI chondrites (1400 to 2400 ppm; Table S-2 and references therein). Notably, particle C0015 contains more Ne and Ar than CI chondrites ( $^{36}\text{Ar} = 4.33 \pm 2.78 \times 10^{-11}$  mol/g and  $^{20}\text{Ne} = 1.53 \pm 0.03 \times 10^{-11}$  mol/g, on average, in CIs; Broadley *et al.*, 2023 and references therein), but the large noble gas abundance predominantly results from SW irradiation.

Broadley *et al.* (2023) suggested that preferential loss of  $^{15}\text{N}$ -rich soluble organic matter during aqueous alteration on Ryugu's parent body may have resulted in a lower  $\delta^{15}\text{N}$  and lower N concentrations in Ryugu samples than in CIs, without significantly affecting the noble gas budget. According to this scenario, the N heterogeneities between the various small, (sub-)milligram-sized Ryugu samples could be due to variable degrees of aqueous alteration on the initial parent body, prior to the catastrophic break-up that turned Ryugu into a rubble-pile asteroid and led to mixing between more and less altered clasts (Broadley *et al.*, 2023). However, this process alone can neither explain the association of "intermediate"  $\delta^{15}\text{N}$  values with high N abundances in particles C0015 (this study) and A0033 (Nakamura *et al.*, 2022), nor the very low  $\delta^{15}\text{N}$  value and high N content of particle C0082 (Nakamura *et al.*, 2022) (Fig. 3). Significant heterogeneities of the N distribution must, therefore, exist at the scale of (sub-)milligram-sized Ryugu regolith particles, as also demonstrated by *in situ* secondary ion mass spectrometry analyses of N (Nakamura *et al.*, 2022).

CI-type material is inferred to have played an important role in supplying volatile elements (*e.g.*, H, C, N) to Earth's surface (Piani *et al.*, 2020). However, since thermal metamorphism and accretionary heating could have resulted in extensive devolatilisation and loss of  $^{15}\text{N}$ -rich components (Alexander *et al.*, 1998; Alexander *et al.*, 2007; Pearson *et al.*, 2006; Grewal, 2022), assessing the evolution of the N content and isotopic composition during progressive heating is key for understanding the contribution of CIs to Earth's volatile budget. The numerous heating steps performed on Ryugu particle C0015 provide key information on N abundance and isotopic variations induced by thermal processing, especially since Ryugu never experienced temperatures  $>150$  °C (based on analyses of soluble organic



**Figure 4** Evolution of the residual N isotopic composition ( $\delta^{15}\text{N}_{\text{residual}}$ ) of particle C0015 as a function of the fraction of N extracted by progressive heating.

matter in Ryugu; Naraoka *et al.*, 2023). Figure 4 illustrates that the "bulk"  $\delta^{15}\text{N}$  value of particle C0015 was not significantly modified during the first 60 heating steps (*i.e.*  $\delta^{15}\text{N} = 22.05 \pm 0.22$  ‰ to  $24.47 \pm 0.17$  ‰ for 1 to 7 % of the maximum laser power), although up to  $\sim 73$  % of the total initial N content was lost. As heating proceeded and particle C0015 started to melt at step #80, the  $\delta^{15}\text{N}$  decreased to  $4.47 \pm 0.47$  ‰ (with concomitant loss of  $\sim 99$  % N). This observation confirms that Ryugu samples contain a refractory N component with an isotopic composition similar to Earth's atmosphere (Hashizume *et al.*, 2024) whose proportion of the total amount of N, however, is very small.

Overall, N-rich material collected at Ryugu's surface is predominantly characterised by an isotopic signature that is comparable to, or only slightly lighter than, that of CI chondrites ( $\delta^{15}\text{N} \geq +18$  ‰; Fig. 3; Table S-2). Furthermore, our new data demonstrate that low temperatures can result in extensive N loss from CI-type material, without significantly affecting the bulk N isotope composition. Only high temperatures result in loss of a  $^{15}\text{N}$ -rich component and a notable decrease of the bulk  $\delta^{15}\text{N}$  value by  $\sim 20$  ‰. Consequently, and despite potential losses of thermally labile N-bearing phases, N isotopes remain a reliable and powerful tool for tracing contributions from inner and outer Solar System sources, and they imply that CI- or Ryugu-type material can be ruled out as the major source of N in Earth's mantle.

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## Additional Information

Supplementary Information accompanies this letter at <https://www.geochemicalperspectivesletters.org/article2431>.



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